

The Imbert Formation
of northern
Hispaniola

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The Imbert Formation of northern Hispaniola: a tectono-sedimentary record of arc-continent collision and ophiolite emplacement in the northern Caribbean subduction-accretionary prism

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Received: 11 May 2015 – Accepted: 25 May 2015 – Published: 26 June 2015

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Published by Copernicus Publications on behalf of the European Geosciences Union.

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Abstract

In northern Hispaniola, the Imbert Formation (Fm) has been interpreted as an orogenic “mélange” originally deposited as trench-fill sediments, an accretionary (subduction) complex formed above a SW-dipping subduction zone, or the sedimentary result of the early oblique collision of the Caribbean plate with the Bahama Platform in the middle Eocene. However, new stratigraphical, structural, geochemical and geochronological data from northern Hispaniola indicate that the Imbert Fm constitutes a coarsening-upward stratigraphic sequence that records the transition of the sedimentation from a pre-collisional forearc to a syn-collisional piggy-back basin. This piggy-back basin was transported on top of the Puerto Plata ophiolitic complex slab and structurally underlying accreted units of the Rio San Juan complex, as it was emplaced onto the North America continental margin units.

The Imbert Fm unconformably overlies different structural levels of the Caribbean subduction-accretionary prism, including a supra-subduction zone ophiolite, and consists of three laterally discontinuous units that record the exhumation of the underlying basement. The distal turbiditic lower unit includes the latest volcanic activity of the Caribbean island arc; the more proximal turbiditic intermediate unit is moderately affected by syn-sedimentary faulting; and the upper unit is a (caotic) olistostromic unit, composed of serpentinite-rich polymictic breccias, conglomerates and sandstones, strongly deformed by syn-sedimentary faulting, slumping and sliding processes. The Imbert Fm is followed by subsidence and turbiditic deposition of the overlying El Mamey Group.

The $^{40}\text{Ar}/^{39}\text{Ar}$ plagioclase plateau ages obtained in gabbroic rocks from the Puerto Plata ophiolitic complex indicate its exhumation at $\sim 45\text{--}40$ Ma (lower-to-middle Eocene), contemporaneously to the sedimentation of the overlying Imbert Fm. These cooling ages imply the uplift to the surface and submarine erosion of the complex to be the source of the ophiolitic fragments in the Imbert Fm, during of shortly after the emplacement of the intra-oceanic Caribbean island-arc onto the continental margin.

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1 Introduction

Intra-oceanic arc systems develop as a result of subduction initiation within oceanic lithosphere and subsequent plate convergence. In this tectonic context, forearc basins form on the upper plate between the arc volcanic front and the outer-forearc high (Dickinson, 1995). They differ of the outboard trench and trench-slope basins, which form along the boundary between the two convergent plates in response to flexure of the subducting slab and extension of the upper plate, often driven by basal tectonic erosion (Draut and Clift, 2012). The distribution and nature of sedimentary sequences in a forearc basin is controlled by uplift or subsidence and by the composition of the rocks eroded from the neighbouring volcanic arc, subduction complex and continental margin (Dickinson, 1995).

Chaotic rock assemblages, often referred in a descriptive non-genetic sense to a *mélange* (Cowan, 1985), are commonly found in forearc and trench environments and are attributed either to tectonic, sedimentary or diapiric processes, as well as their mutual interplay and superposition (Harris et al., 2009; Festa et al., 2010; Pini et al., 2012). They are often directly linked to the subduction zone processes. However, the genesis of any chaotic rock assemblage must be deduced from both a detailed analyses of its lithological content, the source of its components, and the kinematic interpretation of its fabrics, as well as independent evidences concerning its paleotectonic setting. A particular chaotic rock assemblage is the olistostrome, which is a sedimentary body derived from diverse types of gravity mass movements, such as block slides, debris avalanches, debris flows, and hyper-concentrated turbiditic currents (Lucende and Pini, 2003). These “sedimentary *mélanges*” are compatible with the classic principles of stratigraphic superposition, whereas other *mélange* occurrences in nature do not follow these principles because they are bounded by tectonic contacts. Useful criteria to discriminate between chaotic units of sedimentary and tectonic origin are the presence or absence of soft sediment deformation, the source of the rock fragments, the nature of the fine-grained matrix, and the tectono-stratigraphic position within the accretionary

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prism (e.g., Festa et al., 2010). The internal structure and the sedimentary record preserved in mélanges can provide a better understanding of the tectono-stratigraphic evolution of the orogen in which they occur (e.g., Alonso et al., 2006, 2015).

The Greater Antilles orogenic belt results from the Late Cretaceous-Eocene convergence and final collision between the Caribbean island-arc and the North America continental margin, which led the closure of the intermediate proto-Caribbean oceanic domain (Escuder-Virueite et al., 2011c, 2013b; García-Casco et al., 2008; Laó Dávila et al., 2012; Mann et al., 1991; Pindell and Kennan, 2009; Stanek et al., 2009). As a result, several ophiolitic massifs were emplaced in the collisional zone, which are particularly well exposed in Cuba (Lewis et al., 2006; Marchesi et al., 2006). In northern Hispaniola, the Puerto Plata ophiolitic complex forms part of subduction-accretionary prism of the Greater Antilles orogenic belt, which is constituted here by a series of accreted ophiolites, ophiolitic mélanges, intra-oceanic volcanic arcs and fragments of the southern margin of the North America continent (Escuder-Virueite et al., 2011a, b; 2013a, 2014).

In the Cordillera Septentrional (Fig. 1), remnants of a Cenozoic pre- to syn-collision forearc sedimentary basin edified over the orogenic prism are constituted by the Imbert Formation (Fm) and the > 5 km-thick overlying turbiditic sequence of the El Mamey Group (de Zoeten and Mann, 1991). Unlike the well-studied El Mamey Group (de Zoeten and Mann, 1991, 1999), there exist various interpretations of the origin and meaning of the Imbert Fm, as well as their spatial and temporal relationships with other sedimentary units. The Imbert Fm has been interpreted as a Paleocene-lower Eocene, deep-marine sedimentary unit deposited in the outer forearc-trench setting, presumably over the underlying Puerto Plata complex (Nagle, 1979; Pindell and Draper, 1991), an orogenic mélange originally deposited as trench-fill sediments (Pindell, 1985), an accretionary (subduction) complex formed above a SW-dipping subduction zone (Bowin and Nagle, 1982; Bourgois et al., 1982), and a sedimentary record of the cessation of subduction-related magmatism as result of the early oblique collision

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of the Caribbean plate with the Bahama Platform in the middle Eocene (Draper et al., 1994; de Zoeten and Mann, 1999; Hernáiz-Huerta et al., 2012).

In this contribution, we present detailed maps, stratigraphic columns, whole-rock geochemical data of volcanic rocks, and data of syn-sedimentary deformation structures of the Imbert Fm, as well as Ar–Ar cooling ages of its basement. Our main objective is to document the stratigraphy and structure of the Imbert Fm, as well as its relationships with igneous and metamorphic basement units belonging to the northern Caribbean subduction-accretionary prism. We will show that the Imbert Fm is not a subduction mélangé but represents a coarsening-upward sedimentary sequence with an olistostromal upper part, which unconformably overlies the supra-subduction zone (SSZ) ophiolite of the Puerto Plata complex and the collisional suture zone of the Río San Juan complex. Our observations and data suggest that the rocks of the Imbert Fm: (1) contain the latest volcanic activity of the Caribbean island-arc, (2) record a change in the morphology of the forearc that was induced by uplift during the arc-continent collision and ophiolite emplacement, (3) include serpentinite-rich, chaotic rock assemblages deposited in a basin transported on top of the ophiolite, and (4) were followed by a syn-orogenic turbiditic sedimentation in a subsequent phase of regional subsidence (i.e., the El Mamey Group).

2 Geological framework

Located on the northern margin of the Caribbean plate, the Island of Hispaniola (Fig. 1) is a tectonic collage produced by the oblique convergence to final collision of the Caribbean island-arc/back-arc system with the North American plate, which began in the Lower Cretaceous (Mann et al., 1991; Draper et al., 1994). The presence of high-P ophiolitic mélanges in northern Hispaniola indicates that an intermediate proto-Caribbean oceanic basin was consumed by SW-directed subduction (Draper and Nagle, 1991; Saumur et al., 2010; Escuder-Virueite et al., 2011a, c). The arc-related rocks of the Caribbean upper plate have ages that span the Aptian-Maastrichtian interval

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disrupted crust and mantle sections of Caribbean oceanic lithosphere, which records a Cretaceous complex history of extreme crustal thinning and related SSZ magmatism (Escuder-Viruete et al., 2014).

The ophiolitic basement is overlain by the Paleocene-lower Eocene > 500 m-thick section of the Imbert Fm (Nagle, 1979), which is composed of fine-grained turbidites interbedded with white and turquoise, very fine-grained tuffs, pelagic sediments, rare radiolarian cherts and basaltic sills. Pindell (1985) extends the unit including turbiditic beds, or debris flows, of sandstone and conglomerate that contain angular to sub-rounded lithic fragments up to 10 cm of serpentinites, as well as metamorphic, volcanic and green siliceous rocks. Pindell and Draper (1991) describe in detail the Imbert Fm and propose that it is coeval or slightly older than an informal unit composed of serpentinitic-rich breccias, and the lower-to-middle Eocene shallow-water limestones of the La Isla Fm, which typically contains clasts of similar serpentinites. The serpentinitic breccias are recently included at the base, and interbedded with, the lower stratigraphic section of the San Marcos Fm (Hernáiz-Huerta et al., 2012), which is a mud-matrix mélange. However, this interpretation disagree with the presence of middle Eocene to middle Miocene faunas in sedimentary blocks included in the San Marcos Fm (Bourgeois et al., 1982; Hernáiz-Huerta, 2010; Monthel, 2010), as well as the Miocene age of the mudstone matrix (Monthel, 2010; Suárez-Rodríguez et al., 2013, 2015).

For Pindell and Draper (1991) and Hernáiz-Huerta et al. (2012), the Imbert Fm records the early collision of the Caribbean island arc with the Bahamas Platform. This is consistent with the middle to upper Eocene folding and uplift tectonic event described in the Septentrional Cordillera by Mann et al. (1991) and De Zoeten and Mann (1999), as well as in Puerto Rico (Láo Dávila et al., 2012), Haiti and southwestern Cuba (Calais et al., 1992). This tectonic event records the termination of Caribbean arc-related activity and the formed structures are older than the unconformably sedimentary base of the El Mamey Group. This Group is composed by the upper Eocene to lower Miocene Altamira, La Toca and Luperón Fms, which comprise 1500 m of cal-

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careous mudstones and siltstones, sandstones and conglomerates (Nagle, 1979; De Zoeten and Mann, 1999). The cover culminates with the subhorizontal middle Miocene to Pleistocene rocks of La Jaiba and Villa Trina Fms.

2.2 The Río San Juan complex

5 The Río San Juan complex (RSJC; Fig. 2) consists of three elements (Draper and Nagle, 1991; Escuder-Virueite et al., 2013a, b): (1) a core of Mesozoic igneous and metamorphic rocks, (2) a group of Paleocene-lower Eocene to middle Miocene olistostromic, turbiditic and siliciclastic rocks (Imbert, La Toca and La Piragua Fms) that unconformably overlie the core at its periphery, and (3) a subhorizontal cover of upper
10 Miocene to Pleistocene limestones (Villa Trina Fm). Recent work by Escuder-Virueite et al. (20113a) indicates that the large-scale internal structure of the complex consists of an imbricate stack of metamorphic rocks derived from both the Caribbean island arc and the proto-Caribbean lithosphere. In structural ascending order, the major tectonic nappes/units are: the Gaspar Hernández peridotite, the Jagua Clara serpentinite-matrix mélange, and the Cuaba and Morrito units.

The Gaspar Hernández unit is composed of typical abyssal serpentinitized peridotites (Saumur et al., 2010) and minor mafic sills with N-MORB geochemical signatures, suggesting that it is a fragment of the proto-Caribbean oceanic lithosphere (Escuder-Virueite et al., 2011c). The Jagua Clara mélange consists mainly of schistose and sheared, antigorite serpentinite. Blocks-in-matrix eclogites and blueschists
20 originated from N-MORB and IAT type magmas. The retrograde exhumation of the mafic blocks and their re-equilibration to blueschist and greenschist-facies took place by Campanian-Maastrichtian time (Escuder-Virueite et al., 2013b). The Cuaba unit is mainly composed of mafic lithologies metamorphosed to amphibolite and eclogite-facies conditions. They originated from IAT and calc-alkaline type protoliths (Draper and Nagle, 1991; Escuder-Virueite, 2009). The Morrito unit is a nappe of metavolcanic
25 rocks (Draper and Nagle, 1991), which have boninite and IAT type compositions. These protoliths are compositionally similar to the Lower Cretaceous Caribbean island-arc

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rocks (Escuder-Viruete et al., 2011c). The basal thrust zone juxtaposes the Morrito unit northward onto the Jagua Clara mélange. This juxtaposition took place in the latest Maastrichtian to Paleocene, at the onset of the arc-continent collision (Escuder-Viruete et al., 2013b). The Morrito basal thrust has been interpreted as the suture zone between the upper Caribbean island-arc terranes and the lower paleomargin of North America (Escuder-Viruete et al., 2011c).

The unconformably overlying olistostromic rocks of the Imbert Fm are mainly composed of poorly sorted angular breccias and large blocks, with subordinated sandstones, mudstones and rare white and turquoise tuffs (Draper and Nagle, 1991; Escuder-Viruete, 2009). The clasts of the breccias are mainly of serpentinite from the underlying Gaspar Hernández peridotites, but also contain clasts of Jagua Clara and Morrito metamorphic rocks. The source area for the breccias and olistostromes was probably a fault-controlled escarpment, which indicates the exhumation of the subduction-accretionary complex to submarine exposure before the Paleocene-lower Eocene (Escuder-Viruete, 2009). The basal olistostrome is overlain by the middle Eocene to Oligocene La Toca Fm, which is > 1 km-thick and contains mid- to thin-bedded sandstones, siltstones and marls, with rare intercalations of pebbly conglomerates and calcarenites. The sediments of the La Toca Fm are thought to have accumulated in the middle/distal area of a syn-orogenic turbiditic basin. Turbiditic sedimentation is succeeded by alluvial-fan conglomerates of the lower/middle Miocene Piragua Fm, and the upper Miocene-Pliocene marls and reefal limestones of the Villa Trina Fm.

2.3 Tertiary deformation history of the Cordillera Septentrional

De Zoeten and Mann (1999) distinguished three sedimentary intervals in the Tertiary evolution of the Cordillera Septentrional, separated by distinctive regional-scale tectonic events. The first interval includes the Paleocene to lower Eocene deposition of hemipelagic, fine-grained turbidites (Los Hidalgos and Imbert Fm). Sedimentation was terminated by a folding and uplift event, which is thought to be related to early attempted subduction of the Bahama Platform beneath the Caribbean island-

arc. The second interval contains the upper Eocene to lower Miocene deposition of deep-marine, siliciclastic turbidites (Altamira, Las Lavas, and La Toca Fm). Siliciclastic sedimentation was terminated by a folding and uplift event, which is thought to be associated with transpressional strike-slip faulting related to North America-Caribbean plate motion. The third interval includes the upper Miocene to lower Pliocene deposition of shallow-marine limestones (Villa Trina Fm). Carbonate sedimentation was terminated by a folding and uplift event related to the current pattern of restraining bend tectonics (Mann et al., 2002).

3 Stratigraphy of the Imbert Fm

The stratigraphic relationships between the Imbert Fm and its igneous and metamorphic basement are exposed in the PPC and RSJC, which are illustrated with eight stratigraphic sections (Figs. 3 and 4). Three informal stratigraphic units characterize the Imbert Fm: a fine-grained sandstone-mudstone lower IM1 unit, interbedded with fine-grained volcanic rocks (tuffs) and intruded by mafic sills; a coarse-grained sandstone-mudstone intermediate IM2 unit; and an olistostromic and heterogeneous clastic upper IM3 unit, composed of polymictic breccias, conglomerates and sandstones. In the PPC, these units are unconformably overlain by a chaotic pebbly mudstone-mudstone unit, the San Marcos Fm, characterized by a block-in-matrix fabric and containing clasts of rocks similar to, and presumably derived from, all Imbert Fm units, as well as the underlying ophiolitic basement. In the RSJC, the IM1 and IM2 units are absent, and IM3 unit directly overlies the metamorphic basement. The rocks of the IM3 unit are overlain by the turbiditic sequence of the La Toca Fm, which contain in the lowermost part calcarenites with foraminifera of middle-to-upper Eocene age (Escuder-Viruete, 2009). During the fieldwork, distinctive sedimentary facies associations have been used, all of which are detailed described in the Supplement A.

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3.1 El Puerto section

This section corresponds to the type section of the lower IM1 unit and illustrates the gradual change to the intermediate IM2 unit. The section lies along the northern flank of El Puerto syncline (A–A'; Fig. 2), where the bedding trends N110–130° E and dips 14–36° toward the south. The substratum of the section does not clearly outcrop. Three facies assemblages characterize the IM1 unit: thin-bedded, medium-to-fine-grained sandstone (Sm); thin-bedded alternance of fine-grained sandstone, siltstone and laminated mudstone (Sf); and fine-grained tuff with alternating fine-grained tuffaceous sandstone-mudstone (Tv; Fig. 5h). The lower part of the section is ~ 500 m-thick (from 0 to 480 m in Fig. 3) and lack coarse-grained rocks. It is characterized by interbedded light brown thin-bedded sandstone, dark mudstone and green parallel-laminated siltstone, in beds up to 1 m-thick (Sf). In these hemipelagic facies Monthel (2010) found planktonic foraminifera of the lower Eocene. White, cream and turquoise coloured, fine-grained tuffs with laminated internal structure are sporadically intercalated as very thin beds (1–3 cm thick) in the sequence (Tv). The upper part of the section attains ~ 400 m de thickness (from 500 to 900 m in Fig. 3) and is also constituted by fine-grained deposits. However, this upper part is characterized by a progressive higher occurrence of interbedded 5-to-50 cm-thick beds of medium-to-coarse grained sandstone (Sm), which are characterized by size grading and planar laminations, and result a typical turbidite sequence of alternating sandstone-mudstone. The first lenticular breccias with erosive base typical of the intermediate IM2 unit occurs at ~ 880 m and contains 5-to-20 cm-sized subangular serpentinite clasts in a sandy matrix composed of serpentinite, altered pyroxene, red iron oxide, chlorite and red/green mudstone. The El Puerto section is not deformed by syn-sedimentary faulting.

3.2 Rancho Brugal section

This section is 320 m thick, has a 60 m gap in outcrop halfway up, and corresponds to the lower IM1 unit (B–B'; Fig. 2). The lower part of the section exposes ~ 180 m of alter-

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which sporadically intercalates < 1.5 m-thick lenticular beds of clast-supported breccia and angular-to-sub-rounded conglomerates (Gp). Some parts of the section are disrupted and deformed by syn-sedimentary extensional faults (Fig. 5f and g). The upper part of the section is about 240 m thick (i.e., from 340 to 580 m in Fig. 3) and consists of thick and very thick beds of breccia and conglomerate (Gm and Gp), which characterize the IM3 unit (Fig. 5e). Breccias are often composed of amalgamated lenticular beds with erosive base, where individual beds are characterized by centimeter to meter sized fragments in a fine-grained matrix of sandstone and mudstone. Fragments include, in decreasing order of abundance, serpentinite, peridotite, red/green volcanic rocks, laminated sandstone, leucogabbro, (rare) white limestone and red/green mudstone. Large fragments (up to 1 m) of mudstone are found locally within the breccia and are interpreted as rip-up clasts. The breccias are overlain by, and interbedded with, sequences of 0.1-to-0.5 m-thick beds of alternating coarse-grained sandstone and mudstone (Sc). This upper part is strongly deformed by syn-sedimentary extensional faults.

3.4 Loma Seboruco section

This section is located in the eastern Bahía de Maimón (D–D'; Fig. 2) and is constituted by the upper IM3 unit. Although the lower stratigraphic contact does not outcrop, the section is presumably disposed over the Puerto Plata ophiolitic complex. The bedding trends 18–32° E and dips 20–36° toward the east. The section can be subdivided into two parts, where facies show a coarsening-upward evolution. The lower part exposes ~ 30 m of poorly sorted, coarse-grained sandstones (Sc and Sm), in beds of 0.1-to-0.6 m-thick. Sandstones are rich in volcanic lithic grains and detrital serpentinite grains (Fig. 6b and c). These rocks are deformed by syn-sedimentary faulting. The upper part attains ~ 100 m (from 30 to 130 m in Fig. 3) and is characterized by massive, matrix- and clast-supported polymictic breccias and conglomerates (Gm and Gp). They are often chaotic or form a heterogeneous assemblage of clasts in a serpentinite-rich, muddy matrix. The beds have a thick to very-thick thickness and a massive internal structure. Clasts are angular to sub-rounded, of sand to cobble grain-size, and are of ophiolitic

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and sedimentary nature (Fig. 5b and d). Ophiolitic clasts are mainly composed of variably serpentinized peridotite, with subordinate layered gabbro, leucogabbro, dolerite, basalt, and tonalite; sedimentary clasts correspond to typical lithologies of IM1 and IM2 underlying units, as graded sandstone, fine-grained tuff, tuffaceous siltstone and mudstone, as well as rare white limestone. The matrix of the breccia is a serpentinite-rich, medium-to-coarse-grained sandstone. These breccias are unconformably overlain by the sheared pebbly mudstones of the San Marcos Fm.

3.5 Gaspar Hernández section

This section is ~ 240 m-thick and lies along several carries located in the northwestern end of the Río San Juan complex (E–E'; Fig. 2). The bedding trends N100–126° E and dips 10 to 36° toward the southwest. The sedimentary sequence corresponds to the upper IM3 unit and overlies a basement made up of the Gaspar Hernández serpentinites. Sedimentary facies show an evolution (Fig. 4, base to top) from matrix-supported polymictic breccia and conglomerate (Gm), to lensoidal serpentinite blocks in a heterogeneous matrix (Os and Ms), and clast-supported polymictic breccia and conglomerate (Gp and Gm). In the lower part of the stratigraphic section (from 30-to-100 m), the matrix-supported breccia contains (~ 60 %) 5-to-20 cm-sized angular to subrounded clasts in a serpentinite-rich, muddy-sandy matrix (Gm; Figs. 5c and 6a). Clasts are mainly composed of serpentinite and serpentinized harzburgite, with subordinated red/green volcanic rocks, laminated sandstone, ferruginized microbreccia, porphyritic lavas, isotropic gabbro and dolerite. The breccia beds are internally massive and often have an amalgamated lenticular geometry with erosive bases. These rocks show evidence of syn-sedimentary extensional deformation. The breccias are overlain by an olistostromic intermediate subsection of lensoidal serpentinite blocks in a heterogeneous and mostly deformed matrix, composed of serpentinite-rich sandstone, pebbly mudstone and shale (Os; Fig. 5a). Block range in size from few centimetres to several meters (up to 20 m). Blocks and clasts are angular and of a similar nature than in lower breccias. Breccias are intercalated in 0.1-to-2 m-thick beds of massive vario-

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colored (red, green and grey) mundstone-mundslate (Ms), which also contain serpen-
tinitic clasts and blocks. This caotic and strongly faulted intermediate section is at least
50 m-thick. The upper part of the section is characterized by massive, clast-supported
polymictic breccia and conglomerate (Gp, Gm). The sequence attains ~ 30 m (from
160-to-190 m in Fig. 4). The beds often have a lenticular geometry, a massive or nor-
mally graded internal structure, and a clast-supported fabric. Clasts are angular and of
sand to pebble grain-size. These rocks show syn-sedimentary deformation structures.

3.6 Magante section

This section is ~ 10 km east of the Gaspar Hernández section (F–F'; Fig. 2), and
also correspond to the IM3 unit, although the basement here consist of the Jagua
Clara serpentinite-matrix mélange. The bedding trends N90–110° E and dips 60–75°
toward the south. The section can be subdivided into two subsections, where facies
show a crudely finning-upward evolution. The lower part exposes ~ 24 m (from 30 to
54 m in Fig. 4) of massive and matrix-supported, muddy-sandy polymictic breccia and
conglomerate, with cm-sized clasts of ophiolitic, metamorphic and sedimentary rocks
(Gm). Ophiolitic clasts are composed of serpentinite, peridotite, and various types of
gabbro, dolerite and volcanic rocks. Metamorphic clasts consist of eclogite, blueschist,
metagabbro and felsic orthogneiss. Sedimentary clasts are typical lithologies of IM1
unit, as brown graded sandstone, white fine-grained tuff and tuffaceous siltstone, lime-
stone and mundstone. The matrix of the breccia is medium-to-coarse-grained sand-
stone, rich in serpentinite grains. In the upper part of the section, breccias have a clast-
supported fabric (Gp) and are overlain by ~ 16 m of well-bedded and normally-graded
coarse- to fine-grained sandstones. The contacts between these sandstone beds are
commonly scoured. Sandstones contain reworked foraminifera of upper Paleocene age
(Escuder-Viruete, 2009).

3.7 Caño Claro and Hicotea sections

The Caño Claro and Hicotea sections are ~ 1 km apart, and about and 3.5 km south and 9 km east of the Magante section (G–G' and H–H'; Fig. 2). In both sections the unconformable contact between the serpentinite-matrix mélangé and the overlying IM3 unit is exposed. In the Caño Claro section, the metamorphic substrate is directly overlain by ~ 18 m of matrix- and clast-supported polymictic breccia and conglomerate (Gm and Gp), characterized by centimetre-to-decimetre-sized fragments in a serpentinite-rich sandy matrix. Clasts are of similar nature that in the Magante section. Breccia and conglomerate beds (up to 2.5 m-thick) have a crudely finning upward grading, with basal erosive surfaces and coarse-to-medium-grained sandstone tops. In the Hicotea section, the mélangé is directly overlain by only 8–10 m of matrix- and clast-supported, muddy-sandy polymictic breccia (Gp). The matrix contains pebble-to-cobble size clasts of metamorphic rocks, such as metagabbro, eclogite and mafic blueschist. As in the Jagua Clara section, the breccias are immediately overlain by a turbiditic sequence of interbedded sandstone and mudstone of the La Toca Fm, with basal calcarenites containing middle and upper Eocene foraminifera.

4 Volcanic rocks of the Imbert Fm and coeval units

4.1 Field relations and petrography

The volcanic rocks present in the Imbert Fm form up to 10 m-thick coherent bodies and well-stratified, fine-grained volcanoclastic rocks, which are interbedded with the fine-grained sandstones and mudstones of the IM1 lower unit (Fig. 3; Escuder-Virueite, 2010). The coherent bodies are composed of brown-to-dark green mafic rocks, forming lava flows, autoclastic breccias and massive intrusive sills (Fig. 6e). Autoclastic breccias are generally composed of monogenetic clasts. These rocks were erupted in a submarine environment and were intruded by rare syn-volcanic, feeder dykes of mi-

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(40–35) indicate that the magmas were extensively fractionated, with exception of one unfractionated sill in which $Mg\# = 60$ (sample 13JE27). These basalts and basaltic andesites have generally low contents in CaO (6.7–9.0 wt. %), for moderately high Al_2O_3 (16.1–16.8 wt. %) and Fe_2O_{3T} (9.6–12.2 wt. %) contents, which increase progressively for decreasing MgO. These trends are tholeiitic and related to the fractionation of olivine plus Cr-spinel, pyroxene, plagioclase and Fe-Ti oxides.

In the MgO vs. TiO_2 diagram of the Fig. 7, volcanic rocks of the Imbert Fm are compared with the volcanic rocks of the Palma Picada complex and the Hbl-diorites of the Curtiembre plutons. They are also compared to the volcanic rocks of the Lower Cretaceous primitive Caribbean island-arc of Central and Eastern Hispaniola, which record a progressive increase in TiO_2 contents from the boninites and depleted low-Ti island-arc tholeiites (IAT) to normal IAT and back-arc basin basalts (BABB; Escuder-Viruete et al., 2008, 2009), as well as representative Pacific IAT and mid-oceanic ridge basalts (MORB) groups. In terms of MgO contents, the studied samples of the Palma Picada complex and the Curtiembre plutons range from moderately to highly fractionated, including relatively low- and mid-Ti compositions. Samples of the Imbert Fm and Curtiembre plutons have similar TiO_2 contents than the BABB of the Rio Verde complex and the mid-Ti tholeiitic suite of the Mariana Arc-Trough system (Gribble et al., 1998), but lower than the compilation of Pacific MORB (PetDB, 2007). TiO_2 contents of the mafic volcanic rocks of the Imbert Fm are higher to those of the IAT, low-Ti IAT and boninites of the Los Ranchos Fm. However, these rocks are highly fractionated and some samples contain abundant phenocrysts, so they do not provide good estimates of liquid compositions. TiO_2 contents in samples of the Palma Picada complex are similar to the IAT and low-Ti IAT groups.

In the Nb/Y vs. Zr/ TiO_2 immobile trace elements plot of the Fig. 8a, all mafic igneous rocks of the Imbert Fm and Palma Picada complex are subalkalic basalt/andesites and andesites, respectively, and plot within the field of the Caribbean island-arc. For these rocks, Ti/V ratios range between 10 and 20, plotting in the subduction-related field (Fig. 8b). Samples of the Palma Picada complex plot in the IAT and Low-Ti IAT fields

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of the Los Ranchos Fm. However, samples of the Imbert Fm and Curtiembre plutons generally plot at higher Ti and V values, which can be related to the accumulation of Fe-Ti phases in these highly fractionated magmas. In a MORB-normalized multi-element plot (Fig. 8), the volcanic rocks of the Imbert Fm have LREE enriched ($[La/Nd]_N = 1.4\text{--}1.8$) and slight HREE depleted ($[Sm/Yb]_N = 0.8\text{--}1.8$) patterns, with low Nb contents (1.7–2.5 ppm). They exhibit positive Ba, U, K and Pb spikes, and negative Nb-Ta (and Ti) anomalies, typical of subduction-related rocks (Pearce and Peate, 1995). These patterns and the values of the trace element ratios $Ti/V \leq 20$, $Zr/Nb < 5$ and $Zr/Nb > 10$, are characteristic of IAT magmas (e.g., Pearce and Peate, 2005). In this sense, their patterns are very similar to the IAT of the Los Ranchos Fm (Escuder-Virueite et al., 2006). The low TiO_2 content and $(Sm/Yb)_N$ ratios suggest that the mantle source of these mafic magmas was depleted and no contained garnet.

In the Palma Picada complex, three compositional groups have been defined (Escuder-Virueite, 2010): normal IAT, low-LREE IAT and calc-alkaline basalts to andesites. All they display a variably LREE enrichment and a flat HREE pattern, with minor Zr-Hf but marked negative Nb-Ta anomalies (Fig. 8), which also are features typical of subduction-related magmas. The LREE enrichment, monitored by the $[La/Nd]_N$ ratio, increase from the low-LREE IAT (1.1–1.7), to the IAT (1.5–2.0) and calc-alkaline (2.4–3.1) groups. Respect to the IAT group, the low-LREE group of the Palma Picada complex contains lower absolute abundances of HREE and a more prominent negative Zr-Hf anomaly. The LREE depletion, low- TiO_2 contents and lower Ti/V values, as well as the lower HREE levels, suggest that the mantle source for this group was more depleted than for the IAT group of the Imbert Fm. The calc-alkaline group of the Palma Picada complex is characterized by a higher LILE and LREE enrichment, as well as higher $(La/Yb)_N$ ratios, suggesting a more “slab” component added to the depleted source (Pearce and Peate, 1995). Samples of the Curtiembre plutons also display a subduction-related trace element pattern and are compositionally similar to the Imbert Fm and the calc-alkaline group of the Palma Picada complex (Fig. 8). The calc-alkaline character of these Hbl-diorites is reinforced by their relatively high $(La/Yb)_N$

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style of deformation. They consist of conjugate sets of extensional shear fractures and small faults, “en-échelon” vein systems, cataclastic fault rocks, local symmetric and asymmetric pinch-and-swell and boudinage structures, as well as very rare folds. Syn-sedimentary folds have metric amplitudes and are disharmonic, close to isoclinal, and asymmetric with vergence toward the NE and SW. Locally NW–SE hinge lines are present. Folds can occur associated with small normal detachments that display a coherent NE and SE-directed sense of normal movement. All these structures are not pervasively developed and characterize the IM2 and, particularly, the IM3 units.

The stress regime prevailing in the depositional setting can be deduced from the syn-sedimentary deformation structures. In the PPC, the debris flow and turbiditic deposits of the Imbert Fm are affected by WNW–ESE to NW–SE and WSW–ENE to W–E trending joints and small faults (Fig. 4b), which generally present a high angle of dip ($> 70^\circ$). In the RSJC, the debris flows are deformed by W–E to WNW–ESE and NNW–SSE to NNE–SSW trending small faults (Fig. 4d), that have a mid- to high angle of dip ($> 30^\circ$). According to the geometry of the fracture system, they can be classified as hybrid (combined tensional and shear) with two joint sets forming an angle $< 60^\circ$, or shear, organized by two joint or small fault sets with an angle about 60° . As new turbiditic strata seal the hybrid joints and the half-graben structures developed by the movement of the faults, these structures are temporally syn-sedimentary (Fig. 5g). However, to know the original geometry of these structures and to deduce the syn-sedimentary stress regime, they must be restored to the horizontal. Unfolding of the bedding at each station made this correction and results are shown in Fig. 9.

In each station, the corrected orientations of shear and hybrid joints are distributed in two conjugate sets, which generally have high angles of dip ($> 60^\circ$). The more frequent trends are WNW–ESE to NW–SE, W–E to WSW–ENE, and SW–NE. The associated small faults are organized in two similar conjugate sets in each station, which present opposite normal sense of movement. The joints and small faults developed in the gabros and peridotites of the PPC also record these main trends. The restored normal sense of fault movement generally has a minor oblique slip component, developing

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graben and semi-graben structures up to several meters wide. The maximum vertical slips range from several centimetres to few meters. In occasions, the fault planes are gently striated, showing slickensides that enable kinematic interpretation. In several stations, the restored geometry of the brittle structures that affect the sedimentary rocks indicates a general subvertical compression and a NNE–SSW to NE–SW trend of subhorizontal extension (Fig. 9). In the PPC, a NW–SE trend of subhorizontal extension is also locally recorded.

More precise information about the syn-sedimentary stress regime can be obtained from the paleostress analysis of the small brittle structures. The paleostress analysis allows to establish the main stress axis direction and the axial ratio (shape) $R = (\sigma_2 - \sigma_3)/(\sigma_1 - \sigma_3)$ of the stress ellipsoid. For this analysis the *FaultKin* software (Marrett and Almendinger, 1990) was used in nine outcrops of the Imbert Fm (see Supplement C). The results were interpreted in combination with other types of data with kinematic significance, such as conjugate shear fractures, tension fractures and calcite-filled tension gashes.

6 Ar–Ar geochronological data

The main objective of Ar–Ar geochronology was to obtain age constraints for the exhumation of the ophiolitic basement of the PPC, source of the rock fragments found in the Imbert Fm. For this goal, three samples were taken: two gabbros from the ophiolitic substrate and other gabbroic clast from the breccias of the overlying IM3 unit. Sample locations are shown in Fig. 2. Analytical procedures and results are reported in the Supplement D. All ages are quoted at the 2σ level of uncertainty. Geologic time scale is from Gradstein et al. (2012).

Sample 10JE54 is a coarse-grained leucogabbro from the PPC, collected in a quarry on the road Puerto Plata–Imbert. Under the microscope, it displays a layered adcumulate igneous texture, defined by variations in modal contents of plagioclase and clinopyroxene, with minor orthopyroxene, olivine and spinel. The obtained plagioclase plateau

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age is 44.0 ± 3.1 Ma (MSWD = 1.2) for seven steps (1–7) and 100.0% of the ^{39}Ar released (Fig. 10a). The inverse isochron age on these seven points is 42.8 ± 5.3 Ma (MSWD = 1.5), with an initial ^{40}Ar - ^{36}Ar intercept at 350 ± 20 . The inverse isochron has high initial ^{40}Ar - ^{36}Ar argon ratios, indicating the presence of trapped argon with a composition differing from that of atmospheric argon. The isochron age is considered as the best age of the sample, which is similar to the plateau age within error. Sample JM9112 is a coarse-grained leucogabbro from the PPC, collected at the same carry that sample 10JE54. It contains plagioclase, clinopyroxene and orthopyroxene, with minor olivine and spinel. The obtained plagioclase plateau age is 35.1 ± 8.1 Ma (MSWD = 4.6) for eight steps (3–10) and 96.0% of the ^{39}Ar released (Fig. 10b). The inverse isochron age is 34 ± 13 Ma (MSWD = 3.8), with an initial ^{40}Ar - ^{36}Ar intercept at 300 ± 18 . As the inverse isochron $^{40}\text{Ar}/^{36}\text{Ar}$ intercepts equivalent to the atmosphere (295.5), the plateau age is the best estimated age for the sample in spite of its high uncertainty. Sample HH9124 is a block of coarse-grained troctolite from the IM3 unit of the Imbert Fm, collected in the Arroyo Seco outcrop at the Cerro de Gran Diablo. It contains igneous plagioclase and olivine, with minor ortho and clinopyroxene, and secondary hornblende. The obtained hornblende plateau age is 55.0 ± 8.1 Ma (MSWD = 0.1) for six steps (2–7) and 69.9% of the ^{39}Ar released (Fig. 10c). The inverse isochron age is 55.5 ± 9.6 Ma (MSWD = 0.1), with an initial ^{40}Ar - ^{36}Ar intercept at 294.1 ± 9.2 . The sample yielded release spectra with plateau and inverse isochron $^{40}\text{Ar}/^{36}\text{Ar}$ intercepts equivalent to the atmosphere (295.5), and therefore the plateau age is the best estimated age for the sample in spite of its high uncertainty.

7 Discussion

7.1 A Caribbean supra-subduction zone setting for the Imbert Fm volcanism

Geochemical data show that the mafic volcanic rocks of the Imbert Fm present typical subduction-related features (e.g., Pearce and Peate, 1995): LILE (large-ion lithophile

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been deformed by syn-sedimentary extensional faulting. These deposits are interpreted as result of high-concentrated turbidity currents in a proximal submarine fan and/or channel-fill setting, affected by post-depositional slumping or sliding processes. The IM3 unit is a variably thick, olistostromic (caotic) unit, composed of serpentinite-rich polymictic breccias, conglomerates and sandstones. The matrix- and clast-supported breccias are mainly interpreted as product of gravity-driven mass-transport processes (non-cohesive debris flows, debris avalanches, slumps and block slides) in a slope setting, or subaqueous fault scarp. The mesoscopic structures of boudinage, pinch-and-swell and slumping are interpreted as result of layer-parallel extension related to submarine sliding. The serpentinite-rich debris flows, typical of the IM3 unit, are indicative of mobilization of hydrated ultramafic rocks, as those of the underlying PPC and RSJC.

Figure 11 also shown that the IM3 unit unconformable overlies different tectonostratigraphic units in the RSJC (serpentinites at the Gaspar Hernández section and a metamorphic mélangé with high-P blocks at the other sections), establishing clearly that its lower contact is an erosional surface, over an already structured subduction-accretionary prism. In the PPC, the lower stratigraphic contact of the Imbert Fm does not outcrop, but clasts in the breccias are of serpentinitized peridotites, gabbroic rocks and basaltic extrusives, typical of the underlying complex, which suggest that the Imbert Fm unconformably overlie the SSZ ophiolite. Therefore, depositional facies, bed-thickness and grain-size in the Imbert Fm define a large-scale, coarsening-upward sequence, developed between two unconformity surfaces (ES1 and ES2 in Fig. 11). However, an intra-formational unconformity between IM2 and IM3 units cannot be rule out. In the PPC sections, distal turbidites and hemipelagic mudstones, radiolarites and coeval arc-related volcanism, pass upwards into a more proximal turbiditic sequence and to ophiolite-derived breccias and conglomerates. Some of the ophiolite-derived conglomerates include sub-rounded serpentinite and basalt clasts suggestive of a high-energy, more shallow-water setting. This evolution indicates sediment accumulation in a progressively shallower, marine paleoenvironment. By other hand, the Imbert Fm

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is affected by syn-sedimentary, small-scale, high-angle faults in the IM2 unit, and by low-angle extensional faults and block slides in the IM3 unit. This evolution suggests an increase of the tectonic instability upward. The shallowing and the increase in tectonic instability are most likely to have been a response to uplift, erosion and/or tectonic denudation of the underlying basement. These processes could be induced by subduction of the North America continental margin and the isostatic rebound of the overlying Caribbean SSZ ophiolite.

This interpretation agrees with the drastic change of the thickness of the Imbert Fm and the resulting wedge-shaped geometry in cross section for the unit. The wedge-shaped sequence thickens to the SW in the PPC and pinches out at the RSJC to the NE. This geometry is the results of the seafloor morphology created by the rising of the NE edge of the basin, or outer forearc, with exhumation and erosion of the ophiolitic basement and other underlying units of the accretionary prism, as well as the reworking of the turbiditic and volcanic rocks of the IM1 unit to form debris flow deposits. The breccias were shed from subaqueous scarps created by the uplift and faulting and filled rugged seafloor topography produced over the emplaced ophiolitic thrust sheet. In this sense, paleocurrent data indicate the existence of a serpentinite source area in the NE.

7.3 Paleostress analysis of the syn-sedimentary deformations

The results of the paleostress analysis of small fault arrays and joint sets that affect the rocks of the Imbert Fm in selected locations establishes an extensional stress ellipsoid for the syn-sedimentary tectonics in all cases (Fig. 9). These stress ellipsoids are characterized by a subvertical σ_1 axis and a subhorizontal, NNE–SSW to NE–SW σ_3 trending axis. The axial ratios of these ellipsoids are generally high (0.5–0.9; see Supplement C), and indicate a pure to horizontal axially symmetric extension ($\sigma_1 \sim \sigma_2$) stress tensor, which is characteristic of pure normal to normal strike-slip deformation. Although most stress ellipsoids show a NE–SW trending σ_3 axis, in some of them (sites 11JE15, 10JE224B and 10JE53), this orientation corresponds to σ_2 probably due to

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an axe permutation or local perturbations of the regional stress field. The restored syn-sedimentary NW–SE trending normal faults and the meter-scale graben structures that affect the youngest breccia and olistostromic successions of the IM3 unit, like the 11JE51 and 11JE252 sites, all indicate a general NE–SW extension. Therefore, an extensional stress-field affects the uppermost levels of the crust during the sedimentation of the Imbert Fm at least in the lower to middle Eocene.

The extensional tectonics results from instability in the forearc basin. The Imbert Fm records a progressive increase in the syn-sedimentary tectonic activity, culminating in the upper IM3 unit, which include meter to decameter-size blocks/olistoliths of variably serpentinized peridotite. These sedimentary relations suggest gravitational collapse associated with submarine sliding as the main disruption mechanism for the generation of the IM3 olistostromic unit, rather than tectonic deformation. After their initial collapse, different types of mass-wasting phenomena, starting from gravity sliding and followed by slumping and debris flows, contributed to the final emplacement of the blocks as downslope deposits. The NNE–SSW to NE–SW trend of subhorizontal extension indicates that the sedimentary basin was most probably oriented WNW–ESE to NW–SE (actual coordinates). The sub-perpendicular trends of paleocurrent and sliding directions (Figs. 3 and 4) are also consistent with this orientation for the basin.

7.4 Age constraints for the Imbert Fm

$^{40}\text{Ar}/^{39}\text{Ar}$ ages reported for the gabbroic rocks of the PPC would be presumed to be amphibole and plagioclase-cooling ages because all of the dates are from plutonic rocks with high crystallization temperatures. Therefore, the $^{40}\text{Ar}/^{39}\text{Ar}$ plateau ages of 44.0 ± 3.1 Ma and 35.1 ± 8.1 Ma obtained in plagioclase from the gabbros indicate cooling of the PPC at ~ 45 – 40 Ma (middle Eocene), due to the assumed closure temperature of plagioclase (200 – 235 °C). In spite of its high uncertainty, the $^{40}\text{Ar}/^{39}\text{Ar}$ plateau age obtained for hornblende in a block from the IM3 unit of the Imbert Fm also indicate a cooling at $T < 500 \pm 50$ °C (closure temperature of amphibole) in the lower-to-middle Eocene.

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fragments means erosion of a proximal source. Chemical data from the harzburgites and associated boninitic cumulate gabbros are consistent with a highly depleted nature for the PPC ophiolite (Escuder-Virueite et al., 2014), which is a typical feature of SSZ-type ophiolites (e.g., Marchesi et al., 2009). These relationships indicate that breccias and olistostromes formed as culmination of regional uplift and exhumation during the middle Eocene; i.e. during or soon after the latest stages of ophiolite emplacement onto the margin, still in a subaqueous setting. This regional uplift is related to an isostatic response induced by the underplating of (low-density) continental material during the arc-continent collision (Fig. 12). These features are consistent with deformation and sedimentation processes that occurred at shallow crustal levels in a piggy-back basin, located at the top of an advancing accretionary wedge (e.g., Schroetter et al., 2006). In this tectonic context, the Imbert Fm records the evolution of a forearc basin into a true piggy-back basin in the upper Paleocene to middle Eocene, as a consequence of arc-continent collision and SSZ ophiolite emplacement.

This interpretation agree with the general absence of continentally derived clasts in the Imbert Fm, as marbles and micaschists of the Samaná metamorphic complex, suggesting that it is being subducted and metamorphosed beneath the accretionary prism (Escuder-Virueite et al., 2011b). The second phase ended with the compressive deformative event in the middle to upper Eocene that forms folds and thrust in the rocks of the Imbert Fm. This event has been related to the oblique collision of the Caribbean island-arc with the Bahama Platform (Draper et al., 1994; De Zoeten and Mann, 1999).

The third phase includes the deposition of the El Mamey Group in the upper Eocene to lower Miocene. The deep-marine turbidite succession of the La Toca, Altamira and San Marcos Fms implies a period of tectonic subsidence following uplift. These units represent the infilling of a tectonically more mature and relatively stable, onlapping forearc basin. This indicates that uplift of the ophiolite and underlying metamorphic units was followed shortly by subsidence. Subsidence in the basin may have been produced by the progressive movement toward the NE of the orogenic wedge onto the continental margin (Fig. 12).

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The Imbert Fm unconformably overlies the PPC and RSJC. This stratigraphic position and up gradation to the overlying El Mamey Group, is not compatible with a genesis for the Imbert Fm as sediments deposited in a trench or inner trench-slope basin, and then incorporated to a subduction complex (Bourgeois et al., 1982; Pindell, 1985; Hernáiz Huerta et al., 2012). In this sense, the relationships between the Imbert Fm and the underlying SSZ ophiolite and overlying flysch deposits are similar to those of the La Picota Fm of Eastern Cuba, the basal Great Valley Supergroup of the coastal California, the Saint-Daniel mélange in southern Quebec, and the olistostromes, mélanges and terrigenous sediments in Albania, overlying the corresponding Caribbean, Franciscan, Taconian (Appalachian) and Tethyan (Mirdita) ophiolites (Cobiera-Reguera, 2009; Hitz and Wakabayashi, 2012; Robertson and Shallo, 2000; Robertson et al., 2012; Schroetter et al., 2006). In the Greater Antilles orogenic belt, the Imbert Fm represents therefore the base of a syn-collisional piggy-back basin developed in a forearc setting during the arc-continent collision.

**The Supplement related to this article is available online at
doi:10.5194/sed-7-1827-2015-supplement.**

Acknowledgements. The authors wish to thank Gren Draper (Miami University) for their comments on the geology of Dominican Republic and Jacques Monthel (BRGM) for their involvement in mapping and sampling the Puerto Plata area. The Dominican Servicio Geológico Nacional is also thanked for collaboration, particularly to Ing. Santiago Muñoz. The research has been funded by the CGL2009-08674/BTE and CGL2012-33669/BTE projects.

Special Issue: “From orogenesis to geoscience in the service of society: the scientific legacy of Prof. Andrés Pérez-Estaún”

Edited by: J. Alvarez-Marron, F. Bastida, E. Rutter, R. Carbonell, and C. Krawczyk

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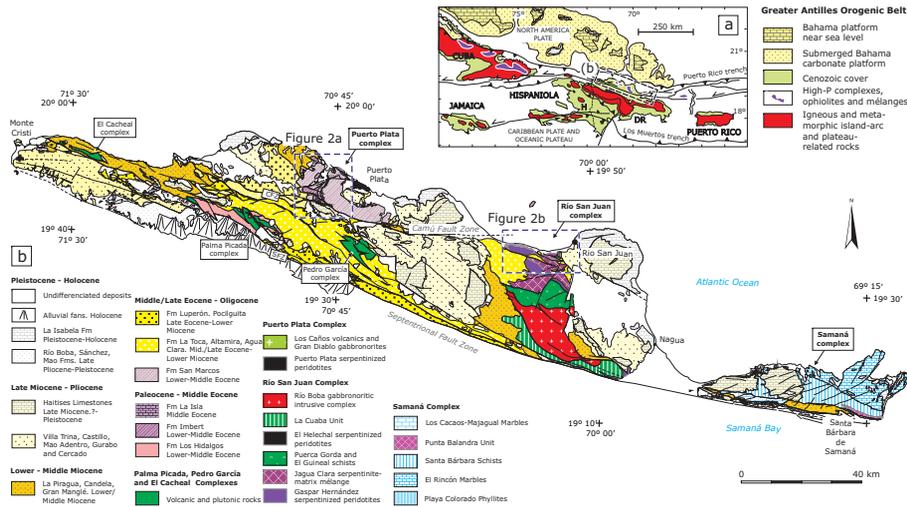


Figure 1. (a) Map of the northeastern Caribbean plate margin. Box shows location of the northern Hispaniola area. DR; Dominican Republic, H; Haiti. (b) Geological map of Septentrional Cordillera and Samaná Peninsula modified from Draper and Lewis (1991), Draper and Nagle (1991) and Escuder-Viruete (2009).

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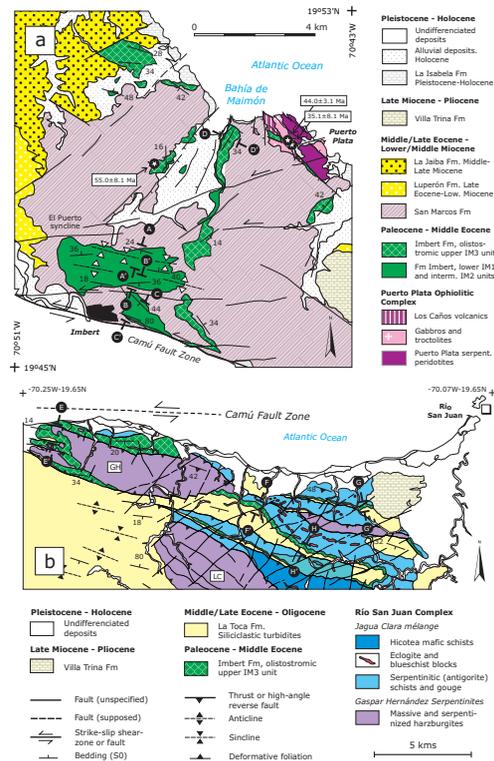


Figure 2. (a) Simplified geological map of Puerto Plata ophiolitic complex modified from Draper and Nagle (1991), Escuder-Viruet et al. (2014) and Suárez-Rodríguez et al. (2015), showing locations of the El Puerto (A–A’), Rancho Brugal (B–B’), Río Obispo (C–C’) and Loma Seboruco (D–D’) sections, the location of the samples dated by $^{40}\text{Ar}/^{39}\text{Ar}$ method, and representative bedding orientations of rocks. **(b)** Simplified geological map of northernmost Río San Juan complex modified from Escuder-Viruet et al. (2013), showing locations of the Gaspar Hernández (E–E’), Magante (F–F’), Caño Claro (G–G’) and Hicotea (H–H’) sections, as well as representative bedding orientations of rocks.

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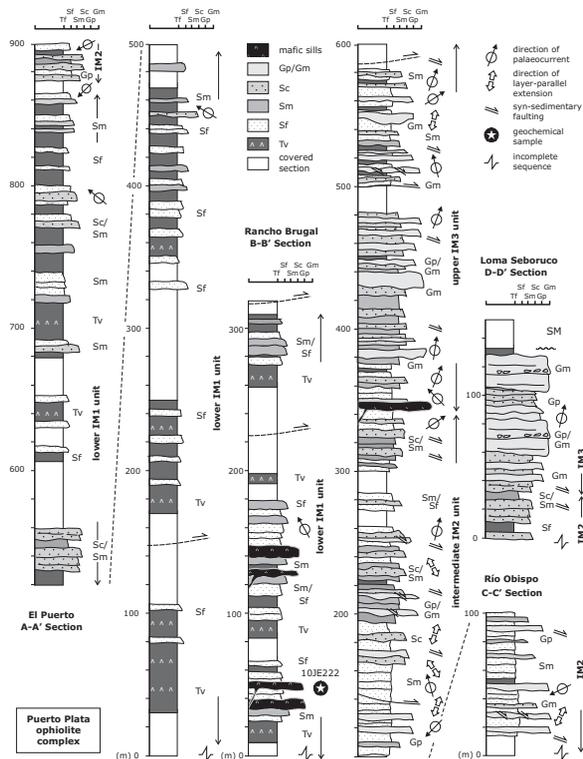


Figure 3. Stratigraphic sections of the Imbert Fm as exposed in the Puerto Plata ophiolitic complex. See text and Supplement A for details and Fig. 2 for location. Sedimentary facies assemblages: Gp, clast-supported, sandy (polymictic) breccia and conglomerate; Gm, massive breccia and matrix-supported, muddy-sandy (polymictic) conglomerate; Sc, medium-to-thick bedded coarse-grained sandstone and microconglomerates; Sm, thin-bedded, medium-to-fine-grained sandstone; Sf, fine-grained sandstone, siltstone and laminated mudstone; Tv, fine-grained tuff with alternating fine-grained tuffaceous sandstone-mudstone.

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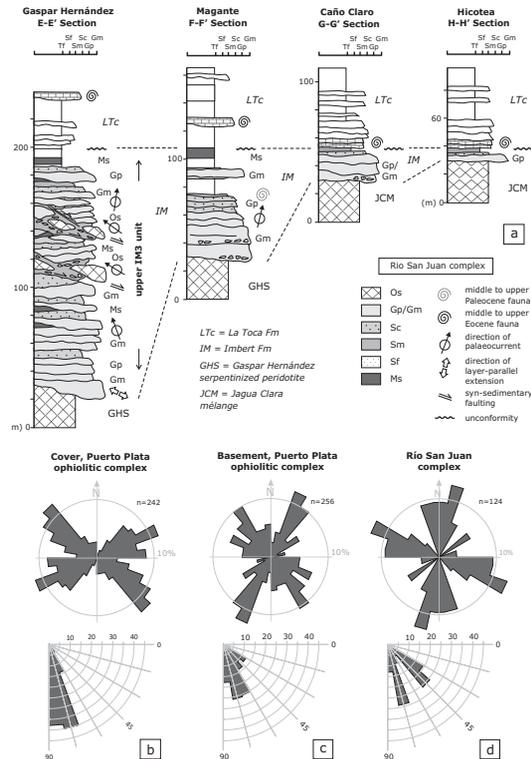


Figure 4. (a) Stratigraphic sections of the Imbert Fm in the Río San Juan complex. Facies assemblages as in Fig. 3, except Os which is composed by lensoidal serpentinite blocks in a heterogeneous and mostly deformed matrix composed of shale, mudstone and sandstone. See text and Supplement A for details and Fig. 2 for location. Trend and dip data of syn-sedimentary joints and small faults affecting (b) the Imbert Fm in the Puerto Plata complex, (c) its ophiolitic basement, and (d) the Imbert Fm in the Río san Juan complex.

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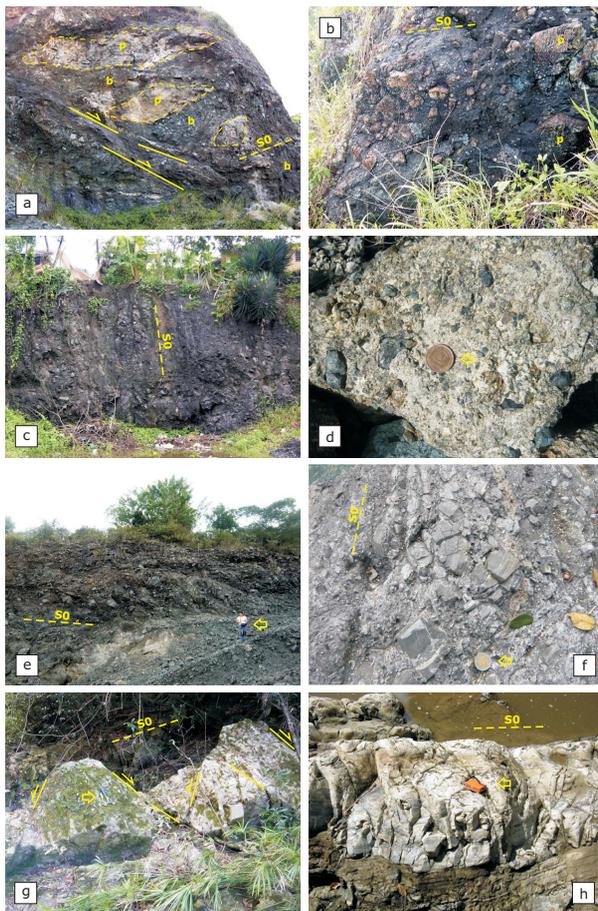
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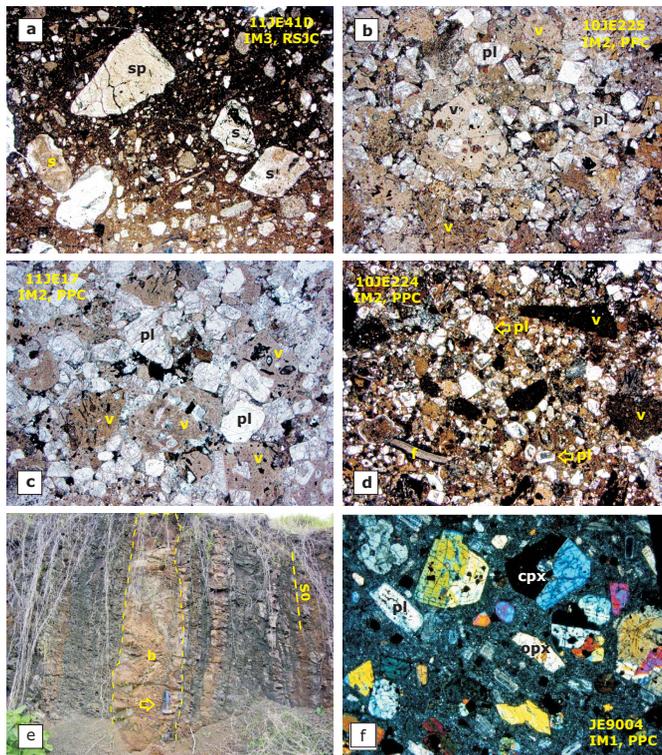


Figure 6. (a–d) Microphotographs (PPL) of the Imbert Fm sandstones and matrix of the breccias consisting of microlitic volcanic rock (v), plagioclase (pl), serpentinized peridotite (sp), serpentinite rock fragments (s), and rare microfossils (f). Volcanic sources of the grains range in composition from basaltic to andesitic. Note the absence of quartz grains. Sandstone grains composition suggest provenance from erosion of volcanic arc and ophiolitic complexes. (e) Basaltic sill intruding an alternance of fine-grained sandstone, siltstone and laminated mudstone of the IM1 unit, Rancho Brugal section. (f) Basalt with microporphyritic textures defined by clinopyroxene (cpx), plagioclase (pl) and orthopyroxene (opx) micro-phenocrysts (XPL).

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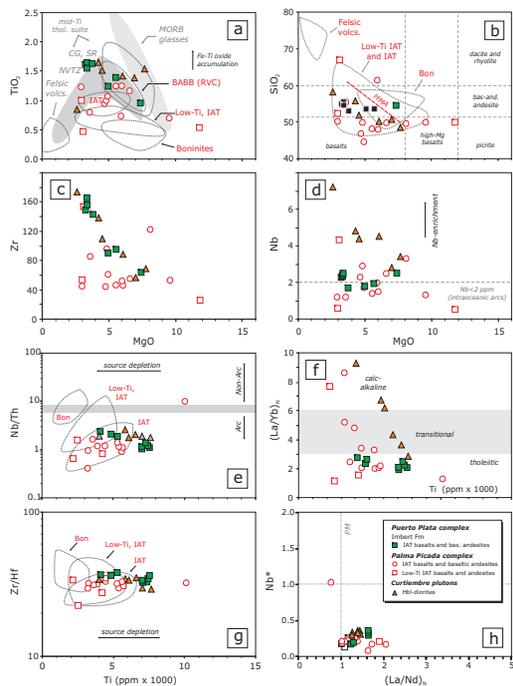


Figure 7. Variation diagrams for rocks from the Imbert Fm, Palma Picada complex and Curtiembe plutons. MgO vs. TiO_2 (**a**), SiO_2 (**b**), Zr (**c**), and Nb (**d**). NVTZ, CG and SR fields are respectively for Northern Volcano-Tectonic Zone, Central Graben and Spreading Ridge fields of the Mariana Arc-Trough system, which are shown for comparison. MORB glasses are from a Pacific compilation. The different geochemical groups of Lower Cretaceous igneous rocks in Hispaniola are: IAT, normal island-arc tholeiites; Low-Ti IAT, Low-Ti and LREE-depleted island-arc tholeiites; Bon, boninites; and felsic volcanic rocks and tonalites. Ti vs. Nb/Th (**e**), $(\text{La}/\text{Yb})_N$ (**f**), and Zr/Hf (**g**), and $(\text{La}/\text{Nd})_N$ vs. Nb^* (**h**) diagrams are for rocks from the same lithological units. See text for explanation.

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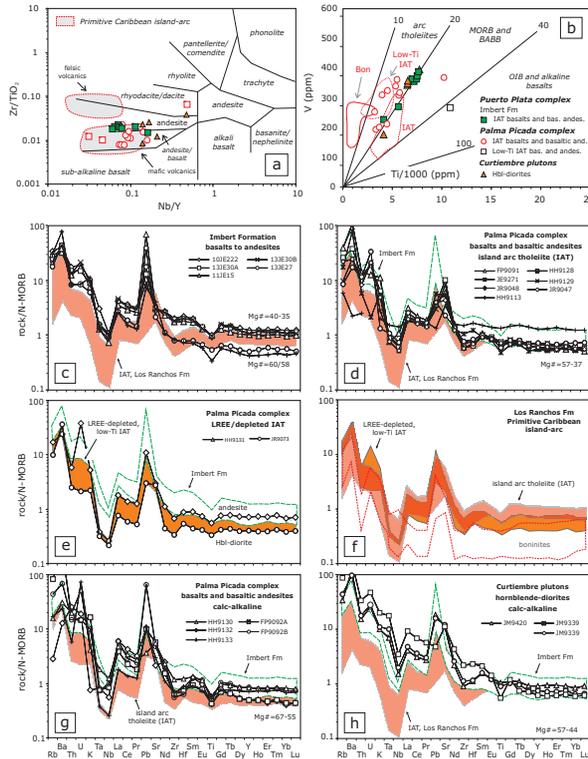


Figure 8. (a) Nb/Y vs. Zr/TiO₂ diagram and (b) Ti–V diagram for rocks from the Imbert Fm, Palma Picada complex and Curtiembre plutons. MORB-normalized extended trace-element plots for volcanic rocks of the (c) Imbert Fm, (d, e and g) Palma Picada complex, and (h) Curtiembre plutons. The diverse geochemical groups of Lower Cretaceous igneous rocks in Hispaniola shown in (f) are: IAT, normal island-arc tholeiites; Low-Ti IAT, LREE-depleted island-arc tholeiites; and Bon, boninites. See text for explanation.

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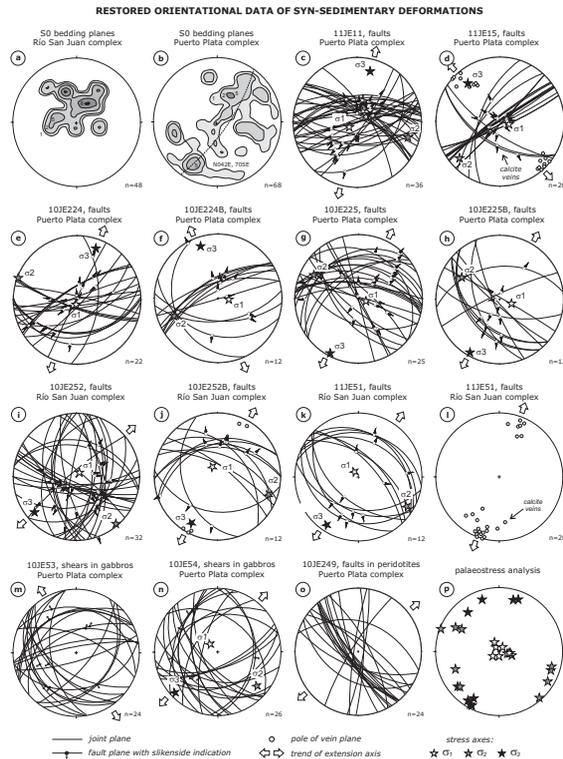


Figure 9. (a and b) Stereoplots of bedding planes (S0) of the Imbert Fm in the Puerto Plata and Río San Juan complexes. (c–o) Stereoplots of fault, joints and calcite veins affecting Imbert Fm deposits from several measurement stations at Puerto Plata and Río San Juan areas. The upper number indicates the station. (p) Stereoplots of stress-axes in all stations. Stereoplots also include the stress-axes obtained from the palaeostress analysis. The white arrows indicate the trend of subhorizontal extension. Equal angle, lower hemisphere projection.

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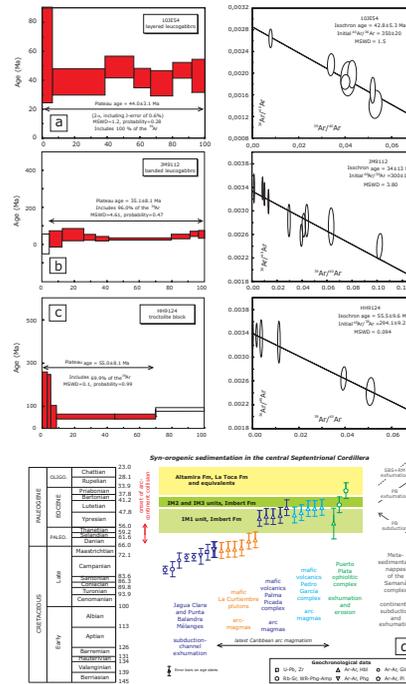


Figure 10. (a–c) $^{40}\text{Ar}/^{39}\text{Ar}$ spectrum and isochron diagrams of plagioclase from leucogabbros of the Puerto Plata ophiolitic complex and of hornblende from a troctolite block in the Imbert Fm. Analytical procedures are described in Escuder-Viruete et al. (2014). A summary of $^{40}\text{Ar}-^{39}\text{Ar}$ incremental heating experiments is in Supplement D. Age uncertainties are 2σ and include uncertainty in monitor age and decay constant. **(d)** Temporal relationships between the exhumation of the Jagua Clara mélangé (end of oceanic subduction), the latest Caribbean arc magmatism, the exhumation of the SSZ ophiolite, and the sedimentation and exhumation of the Samaná nappes (continental subduction). See text for discussion.

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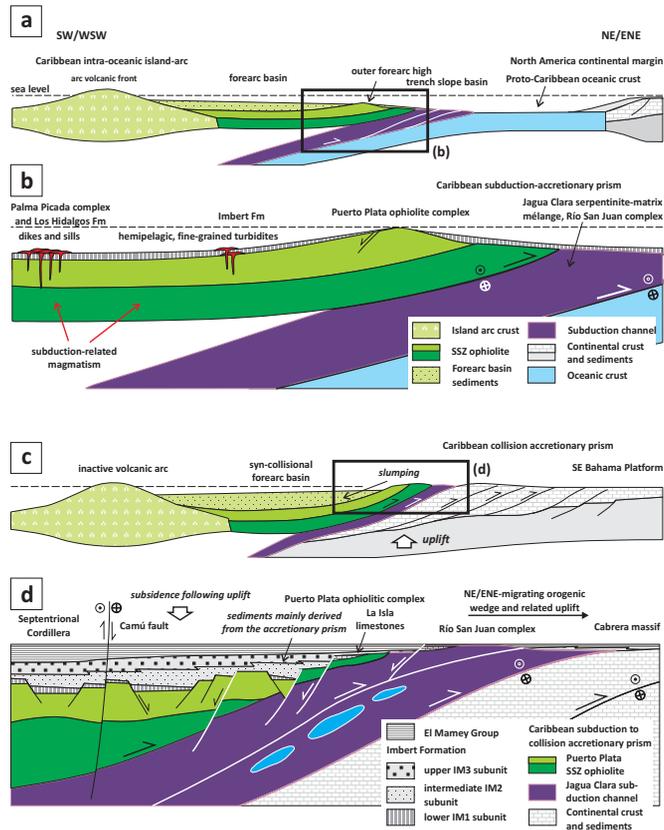


Figure 12. Schematic sedimentary and tectonic evolution of the Imbert Fm and the basal El Mamey Group along the northeastern edge of the pre-collisional forearc (**a** and **b**) and the syn-collisional piggy-back basin (**c** and **d**) of the northern Caribbean subduction-accretionary prism.

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